# Controls of organic and inorganic carbon in randomly selected Boreal lakes in varied catchments

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**Abstract** Organic and inorganic carbon concentrations in lakes and the links to catchment and water quality were studied in variable landscapes using the Finnish Lake Survey data base including 874 randomly selected lakes sampled during autumn overturn. The median total organic carbon (TOC) in these boreal lakes was  $7.8 \text{ mg } l^{-1}$ , the median total inorganic carbon (TIC) 1.6 mg l<sup>-1</sup> and the median partial pressure of CO<sub>2</sub> (pCO<sub>2</sub>) 900 µatm. When the data was divided into subgroups according to land use in the catchment, the proportion of TIC of the total carbon (TC) in lakes was highest (31%) in agricultural areas and lowest (10%) in peatland areas. Elevated TIC concentrations were associated with agricultural land in the catchment, whereas elevated TOC concentrations were observed in lakes with high peatland proportion in the catchment. Two contrasting important sources of CO2 in lakes were identified on the basis of statistical analysis of the data; weathering processes in the catchments and decomposition of organic matter. CO<sub>2</sub> was also strongly associated with total nutrients TN and TP, implying the importance of quality of organic matter and availability of nutrients for the decomposition processes.

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#### Introduction

Organic and inorganic carbon occur in boreal lakes as particulate organic and inorganic carbon (POC, PIC), dissolved organic and inorganic carbon (DOC, DIC) and gaseous forms (free CO2 and CH4). Dissolved inorganic carbon (HCO<sub>3</sub><sup>-</sup>, CO<sub>3</sub><sup>2-</sup>) is related to gaseous carbon (free CO<sub>2</sub>) via carbonate equilibria (Wetzel 1983). The relative proportion of these three inorganic forms of carbon are dependent on pH and to a lesser extent temperature (Stumm and Morgan 1981).

Large rivers have been shown to transport large quantities of inorganic carbon into ocean (e.g. Raymond and Cole 2003; Striegl et al. 2007; Raymond et al. 2008), but the role of total inorganic carbon (TIC) in lakes has received less interest (Dillon and Molot 1997). Bicarbonate (HCO<sub>3</sub><sup>-</sup>) originates mainly from weathering reactions in soil. The chemical weathering of rock material, mainly carbonate and silicate minerals, produces bicarbonate ions (HCO<sub>3</sub><sup>-</sup>) and consumes CO<sub>2</sub> of the atmospheric origin (Gaillardet et al. 1999). The soil CO<sub>2</sub> needed in weathering reactions is mainly derived from respiration of organic matter, but ultimately of the atmospheric origin fixed during photosynthesis (Cole et al. 2007). The resulting bicarbonate ion (HCO<sub>3</sub><sup>-</sup>) is the dominant form of dissolved inorganic carbon in



most lakes and rivers. However, an important source of  $CO_2$  in lakes with longer residence times is also the decomposition of organic matter in situ (Sobek et al. 2003; Kortelainen et al. 2006), either by microbial utilization or photo-oxidation (e.g. Salonen and Vähätalo 1994; Granéli et al. 1996).

Dissolved organic carbon (DOC) is supplied to aquatic ecosystems from both external (allochthonous) and internal (autochthonous) sources. Autochthonous sources include the primary production of algae and macrophytes, whereas allochthonous DOC is derived from the surrounding terrestrial ecosystem. Organic carbon budgets of lakes show that external loading often dominates over internal production (e.g. Canham et al. 2004). However, when production in the littoral zone is added to the carbon budget, at least some clearwater lakes can be net autotrophic (Andersson and Kumblad 2006).

Land use of the catchment affects the quantity and bioavailability of dissolved organic matter (DOM) moving from terrestrial ecosystems into streams (Findlay et al. 2001). The differences in the ultimate origin of DOM combined with differential transformation during transport may result in significantly different behavior of autochthonous versus allochthonous DOM (Findlay and Sinsabaugh 2003). Allochthonous DOM has often experienced biogeochemical reactions before entering the ecosystem. Thus, on average it is less labile than DOM recently produced within the system (Sinsabaugh and Findley 2003). Compared to total organic carbon (TOC), only few studies have described the sources and behavior of TIC in natural waters. A few studies have focused on both organic and inorganic forms of carbon in running waters (e.g. Neal and Hill 1994; Palmer et al. 2001; Dawson et al. 2002, 2004; Johnson et al. 2006; Billett et al. 2007; Striegl et al. 2007; Walvoord and Striegl 2007), but very few in lakes (Dillon and Molot 1997).

This study is based on a randomly sampled database of 874 lakes representing a wide range of size, water quality, catchment land use, and covering a large geographical area, altogether 51% of the total lake area of Finland. Using this extensive data, we identify the concentrations and the relative importance of different forms of carbon (inorganic vs. organic) in lakes surrounded by forests, peatlands and agricultural land, and furthermore, study the potential sources and most important predictors of TOC, TIC and CO<sub>2</sub>.

We hypothesize that in different land use patterns primary drivers controlling the concentrations of TOC and TIC vary. Furthermore, we hypothesize that in complex catchment systems with wide variation in land use, the concentrations of TOC and TIC in lake water are not closely related because of multiple allochthonous sources and several factors regulating TOC, TIC and pCO $_2$  in lakes.

#### Methods

The data set is based on a Nordic lake survey conducted in the autumn of 1995 (Henriksen et al. 1996; Mannio et al. 2000; Rantakari et al. 2004). Altogether 874 lakes were selected from the national lake register using stratified random sampling with unequal sampling fractions, with the requirements that a minimum of 1% of the population of lakes within any county/region was included and that the proportions of lakes in size classes 0.04-0.1, 0.1-1, 1-10 and 10-100 km<sup>2</sup> were 1:1:4:8, respectively. All of the lakes >100 km<sup>2</sup> were included. Larger lakes were emphasized, because lakes >10 km<sup>2</sup> represent over 65% of the total lake area in Finland. One water sample was taken from the middle of the lake during the autumn overturn from depth of 1 m. During autumn circulation a single sample most representatively reflects the water quality of the lake. An extensive seasonal lake survey (Kortelainen et al. 2006) showed that in 1 m samples CO<sub>2</sub> concentrations were only slightly elevated in autumn compared to spring and summer but were significantly lower than in winter samples. Surface water samples of this data also showed low seasonal variability in TOC concentrations as well as in total carbon (TC) concentrations (P. Kortelainen, unpublished). Therefore, autumn sample can be considered as representative for the open water period.

Samples for TIC were taken into airtight glass bottles and placed in coolers while in transit to the laboratories. TIC was measured in the laboratory using infrared spectroscopy. CO<sub>2</sub> concentrations were calculated from measurements of TIC and pH with correction for in situ water temperature (Stumm and Morgan 1981; Butler 1982; Kling et al. 1992). Partial pressure of CO<sub>2</sub> (*p*CO<sub>2</sub>) was calculated from CO<sub>2</sub> concentration using the appropriate Henry's law constant, corrected for temperature and atmospheric



pressure (Plummer and Busenberg 1982). The equilibrium  $p\text{CO}_2$  values in the lake water were calculated with Henry's law, assuming the atmospheric mixing ratio of 360 ppmv for 1995 (IPCC 2001). HCO<sub>3</sub> was calculated as the difference between TIC and CO<sub>2</sub>. The concentrations of  $\text{CO}_3^{2-}$  were assumed to be negligible, since the maximum pH in this data set was 7.9. Similarly, the proportion of CH<sub>4</sub> in these lakes was assumed to be negligible, because when CH<sub>4</sub> in different seasons was measured in the subset of 209 lakes that were selected from Nordic Lake Survey database of 874 lakes, the proportion of CH<sub>4</sub> of TIC in autumn was less than 0.5% (Juutinen et al. 2009).

TOC was analyzed from unfiltered samples by oxidation to CO<sub>2</sub> followed by IR-measurement. Total nitrogen (TN) was analyzed colorimetrically after oxidation to NO<sub>3</sub>-N, the sum of NO<sub>3</sub>-N and NO<sub>2</sub>-N colorimetrically by autoanalyzer after reduction to NO<sub>2</sub>-N, NH<sub>4</sub>-N colorimetrically with hypochlorite and phenol, and organic nitrogen (TON) was calculated as the difference between total and inorganic nitrogen. Total phosphorus (TP) was measured by a colorimetric method after oxidation, phosphate phosphorus (PO<sub>4</sub>–P) by spectrophotometric determination, calcium (Ca) and magnesium (Mg) were determined by flame atomic adsorption/ICP-MS, silicate (SiO<sub>2</sub>) by a colorimetric method or by FIA and sulfate (SO<sub>4</sub>) by ion chromatography (Henriksen et al. 1996; Mannio et al. 2000).

The catchment areas of the lakes were determined from topographic maps. The catchment boundaries were digitalized, and combined with land use data using the Arc View georeferencing software. The land use data is based on satellite images interpretation. Catchment area and lake area were determined, as well as proportions of peatland, (consisting of open peatlands and forests on peatland); forests on mineral soil; agricultural land, (consisting of cropland and pasture); water (consisting of the lake itself and the upstream water bodies); and built-up area in the catchments. The location of the lake in the south-north direction was expressed as latitude (geographical coordinates, minutes and seconds). The catchments were situated between latitudes 60 and 69°N (Fig. 1).

The relationships between lake water TOC, TIC or  $pCO_2$  and catchment characteristics or lake water quality were studied by principal component analysis

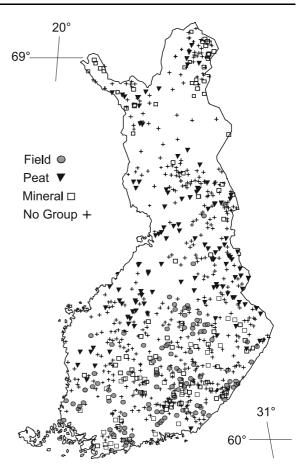


Fig. 1 The locations of the study lakes:  $\blacktriangledown$  subset PEAT (peatland > 35% and agricultural land < 8%), ⊚ subgroup FIELD (agricultural land > 10% and peatland < 12%), □ subgroup MINERAL (forests on mineral soil > 75%, agricultural land < 7% and peatlands < 7%), + the rest of the study lakes (not belonging to any of the subgroups above)

(PCA), correlation and stepwise multiple regression analysis (PROC FACTOR, PROC CORR and PROC REG; SAS Institute 2001). Catchment characteristics (catchment area, lake area, proportions of peatland, forests on mineral soil, agricultural land, water and built-up area in the catchment and latitude) or water quality variables (TOC, TIC, TN, the sum of NO<sub>3</sub>–N and NO<sub>2</sub>–N, NH<sub>4</sub>–N, TON, TP, Fe, Ca, Mg, SiO<sub>2</sub>, SO<sub>4</sub>) were used as independent variables. Most of the concentration and catchment data were transformed into natural logarithms or square roots in order to improve the normality of the distributions. Similar regression models for TOC as a function of catchment characteristics of these 874 randomly selected lakes were presented by Rantakari et al. (2004).



In the regression analysis, cases with the absolute value of studentized residual exceeding three were excluded.

The data was divided into subsets according to different land use in the catchments: (1) a subset with high peatland proportion > 35% and low agricultural land proportion < 8% (PEAT). (2) A subset with high agricultural land proportion > 10% and low peatland proportion < 12% (FIELD). (3) A subset with high proportion of forests on mineral soil > 75% and low proportions of agricultural land < 7% and peatlands < 7% (MINERAL). Cutoffs for the subsets were made using three criteria: (1) a sufficient number of lakes in each subset to enable proper statistical analysis. (2) Approximately equal number of lakes in each subset, because very different number of cases could lead to interpretation difficulties in statistical analysis. Therefore, in each subset cutoffs were chosen to give about 120 lakes. (3) Different medians/

**Table 1** Median characteristics of lakes and catchments in the whole data set, and in the subsets FIELD (agricultural land > 10%, peatland < 12%), PEAT (peatland > 35%,

averages of land use variables between subsets (Table 1).

## Results

Total organic carbon

The size of the lakes ranged from 0.04 to 1, 538 km², the median lake size being 0.22 km². The median TOC in this data set was 7.8 mg l<sup>-1</sup> (range 0.4–34 mg l<sup>-1</sup>) (Table 1). Proportion of TOC of TC was highest in the subset PEAT (90%) and lowest in the subset FIELD (69%) (Fig. 2). The proportion of peatlands correlated positively with lake water TOC in the whole dataset and in the subset PEAT (Table 2). However, in the subsets MINERAL and PEAT agricultural land in the catchment showed stronger correlation with TOC than peatlands, and

agricultural land < 8%) and MINERAL (forests on mineral soil > 75%, agricultural land < 7%, and peatlands < 7%)

n	Whole data set 874	FIELD 120	PEAT 117	MINERAL 119	
Lake area (km²)	ake area (km²) 0.22 (0.04–1,500)		0.12 (0.04-43)	0.090 (0.04–32)	
Catchment area (km²)	4.6 (0.08-52,000)	7.3 (0.16–17,000)	3.4 (0.23-1,300)	1.5 (0.18–370)	
Water (%)	8.8 (0.2–43)	9.6 (0.3–35)	5.2 (1-42)	9.7 (1–24)	
Agricultural area (%)	2.3 (0-53)	19 (11–53)	0.2 (0-7)	0 (0-6.5)	
Forest on mineral soil (%)	65 (10–98)	64 (32–83)	48 (10–61)	85 (76–98)	
Peatland (%)	14 (0–84)	3.0 (0-11)	44 (36–84)	2.9 (0-6)	
Built-up area (%)	0.09 (0-39)	0.60 (0-39)	0.01 (0-0.3)	0.03 (0-13)	
CA/LA	16 (2.1–6,800)	13 (2.7–6,800)	25 (2.2–1,500)	12 (2-640)	
$TOC (mg l^{-1})$	7.8 (0.4–34)	8.1 (0.6–18)	10 (1.3–34)	6.2 (0.4–22)	
TIC (mg $l^{-1}$ )	1.6 (0.1–12)	3.6 (0.1-8.2)	1.1 (0.2–7.3)	1.5 (0.25–12)	
pCO <sub>2</sub> (μatm)	900 (75–5,800)	1,200 (140-4,000)	810 (160–1,300)	930 (260-2,000)	
Water color (Pt mg l <sup>-1</sup> )	60 (3.8–500)	40 (3.8–200)	100 (5–360)	35 (3.8–500)	
pH	6.7 (3.8–7.9)	6.9 (3.8–7.7)	6.4 (4.2–7.5)	6.6 (4.9–7.9)	
Alkalinity (μeq l <sup>-1</sup> )	120 (-170-980)	270 (-170-980) 80 (-80-590)		100 (-20-970)	
TN ( $\mu g l^{-1}$ )	430 (60–2,500)	600 (82–1,800)	460 (150–1,800)	340 (60–790)	
TON ( $\mu g l^{-1}$ )	380 (46–2,100)	500 (69–1,300)	430 (120–1,800)	300 (46–770)	
TOC:TON	25 (5.5–77)	19 (5.7–34)	30 (9.5–77)	24 (5.5–56)	
TP ( $\mu g l^{-1}$ )	13 (1–200)	23 (4–100)	18 (4–88)	9.0 (1-100)	
TOP (TP-PO <sub>4</sub> )( $\mu$ g l <sup>-1</sup> )	12 (1–180)	17 (3–58)	15 (2.5–75)	7.8 (1–55)	
Fe ( $\mu$ g l <sup>-1</sup> )	360 (2.5–5,000)	310 (40–3,500)	760 (20–3,400)	180 (2.5–3,400)	
$Ca (mg l^{-1})$	2.8 (0.2–45)	5.7 (1.7-45)	1.7 (0.2–10)	2.5 (0.4–12)	
$SiO_2 (mg l^{-1})$	2.8 (0.05–21)	2.5 (0.05–17)	2.6 (0.1–18)	2.9 (0.1–10)	



Fig. 2 The relative contribution (%) of each C-species to the total C concentration in the whole dataset and in the subsets PEAT, FIELD and MINERAL (see Fig. 1 for explanation of abbreviations)

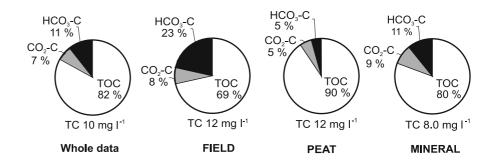


Table 2 Pearson correlation coefficients between log[TOC] and catchment and water quality variables for the whole dataset and for the subsets FIELD, PEAT and MINERAL

See Table 1 for explanation of abbreviations
Significant coefficients
\* P < 0.05, \*\* P < 0.001,
\*\*\* P < 0.0001 are shown,
NS not significant

n	Whole data 814	FIELD 120	PEAT 117	MINERAL 119	
pН	-0.371***	-0.195*	-0.476***	-0.588***	
Peatland (%)	0.340***	NS	0.243**	NS	
Agr. land (%)	0.182***	NS	0.270**	0.246**	
Water (%)	-0.433***	-0.563***	-0.395***	NS	
Built-up (%)	NS	NS	NS	NS	
CA/LA NS		0.398***	0.202*	NS	
Color (Pt mg l <sup>-1</sup> )	0.888***	0.804***	0.904***	0.887***	
TN ( $\mu g l^{-1}$ )	0.744***	0.721***	0.513***	0.762***	
TON ( $\mu g l^{-1}$ )	0.774***	0.711***	0.583***	0.857***	
TP ( $\mu g l^{-1}$ )	0.688***	0.395***	0.573***	0.697***	

furthermore, in the subset FIELD catchment to lake area ratio (CA/LA) had a positive relation with TOC (Table 2; Fig. 3a). Upstream water bodies were negatively related to TOC (Table 2; Fig. 4), as was already shown by Rantakari et al. (2004). Correlation coefficients between TOC and the upstream lake percentage were better in agricultural areas than in peatlands or mineral soils (Table 2).

TOC:TON ratio and water color differed between subsets (Table 1). In the subset PEAT, the median water color and TOC:TON ratio were highest, and significantly lower TOC:TON and color values were recorded in both the FIELD and FOREST subsets, suggesting differences in the quality of lake water organic matter. In PCA analysis, TOC was strongly associated with TN and TON in the subset MIN-ERAL but the connection was less clear in the subsets PEAT and FIELD (Fig. 3a–c).

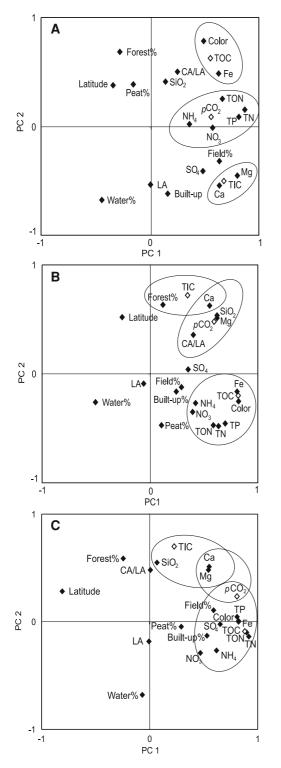
## Total inorganic carbon

The median TIC in this data set was  $1.6 \text{ mg l}^{-1}$  (range  $0.1-12 \text{ mg l}^{-1}$ ) (Table 1), and the average pH

was 6.7 (range 3.8–7.9), indicating that TIC in the lakes consists of bicarbonate and free CO<sub>2</sub>. The proportion of TIC of TC was the highest, 31%, in the subset FIELD. The subsets PEAT and FIELD had the same average TC concentration (12 mg l<sup>-1</sup>), but there were considerable differences in the proportions of inorganic and organic carbon. Organic carbon dominated in both subsets but the proportion of TIC of TC was significantly higher in the subset FIELD (31%) than in the subset PEAT (10%) (Fig. 2).

Bicarbonate dominated alkalinity in these lakes, and consequently alkalinity explained 95% of the variability of TIC in the whole dataset (Fig. 5). As well as alkalinity, TIC was closely related to Ca, Mg and  $SO_4$  (Table 3, Fig. 4). When testing the relationship between TIC and land use in the catchment we found that agricultural land and built-up land in the catchment correlated positively with TIC concentrations, whereas peatlands correlated negatively with TIC (Table 3; Fig. 4). TIC correlated positively with  $pCO_2$  in the whole dataset and in all the subsets (Table 4), but had non-significant correlations with TOC in the whole dataset and in all the subsets (Table 3).





**Fig. 3** The PCA analysis of the water quality and catchment variables in the subsets **a** FIELD, **b** PEAT and **c** MINERAL (see Fig. 1 for explanation of abbreviations)

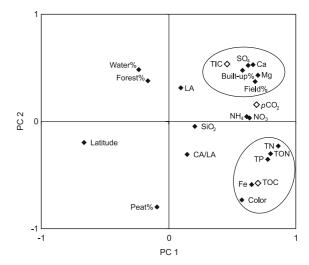
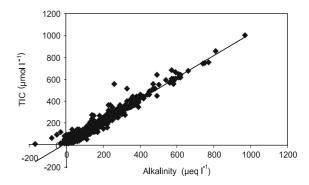


Fig. 4 The PCA analysis of the water quality and catchment variables in the whole dataset



**Fig. 5** The relationship between alkalinity ( $\mu$ eq l<sup>-1</sup>) and TIC content ( $\mu$ mol l<sup>-1</sup>) of the lake water in the whole data set (y = 0.996x + 19.3;  $R^2 = 0.95$ )

Impacts of water quality and land use on CO<sub>2</sub>

Median  $pCO_2$  in the data was 900  $\mu$ atm (range 75–5,800  $\mu$ atm) (Table 1). The highest median  $pCO_2$  was found in the subset FIELD (1,200  $\mu$ atm), in which partial pressure was significantly higher than in the subset PEAT (810  $\mu$ atm). The proportion of  $CO_2$  of TIC varied between subsets, being highest in the subset PEAT (average pH 6.4), in which 53% of the TIC was in the form of  $CO_2$  and lowest in the subset FIELD (average pH 6.9) in which the proportion of  $CO_2$  was only 25% (Fig. 2).

Principal component analysis of the data suggested that there are two major environmental drivers determining  $CO_2$  in the lakes: on the one hand  $CO_2$ 



Table 3 Pearson correlation coefficients between log[TIC] and catchment and water quality variables for the whole dataset and for the subsets FIELD, PEAT and MINERAL

	Whole data	FIELD	PEAT	MINERAL 119	
n	814	120	117		
pН	0.679***	0.658***	0.576***	0.603***	
Peatland (%)	-0.339***	-0.340***	NS	NS	
Agr. land (%)	0.465***	0.519***	NS	0.250**	
Water (%)	NS	NS	NS	NS	
Built-up (%)	0.357***	0.390***	NS	NS	
Forest (%)	NS	-0.433***	0.201*	NS	
Latitude	NS	NS	0.422***	NS	
$TOC (mg l^{-1})$			NS	NS	
TN ( $\mu g \ l^{-1}$ )			NS	NS	
TP ( $\mu g l^{-1}$ )	0.131***	0.468***	NS	NS	
Ca $(mg l^{-1})$	0.810***	0.814***	0.807***	0.766***	
$Mg (mg l^{-1})$	0.774***	0.863***	0.768***	0.705***	
$SO_4 (mg l^{-1})$	0.354***	0.377***	NS	NS	
$SiO_2 (mg l^{-1})$	0.266***	NS	0.445***	0.238**	

See Table 1 for explanation of abbreviations Significant coefficients \*P < 0.05, \*\*P < 0.001, \*\*\*P < 0.0001 are shown, NS not significant

was associated with Ca, Mg and  $SO_4$  concentrations, that are linked to weathering processes in the catchment, and on the other hand  $CO_2$  was associated with TOC, TN, TON and TP concentrations and water color, i.e. lake water organic matter concentrations (Figs. 3, 4). PCA analysis of the subsets showed variation in controls of  $pCO_2$  in different land uses. In the subset PEAT,  $pCO_2$  had a strong relationship with  $SiO_2$ , Mg and Ca, but did not have a close relationship with TOC, or with the different forms of nitrogen or TP. The situation was opposite in the subset FIELD, in which  $pCO_2$  had the closest relationship with TN, NO<sub>3</sub>, NH<sub>4</sub> and TP (Fig. 3a, b).

The stepwise multiple linear regression equations supported the hypothesis of two environmental drivers in regulating  $pCO_2$  by selecting Ca representing the weathering processes and either TOC, TN or TP representing organic matter (Table 5). Additionally, SiO<sub>2</sub> was selected in all equations. In PCA analysis, SiO<sub>2</sub> was not strongly associated with the first two principal components, but mainly with the third principal component, which was characterised by positive loadings for CA/LA and SiO<sub>2</sub> and the negative loading for water in the catchment.

Of the water chemistry variables (besides TIC), Ca had the best correlation with  $pCO_2$  in the whole dataset. In the subsets PEAT and MINERAL, SiO<sub>2</sub> and TOC, respectively, had the best correlation with  $pCO_2$  (Table 4; Fig. 6a). TOC correlated positively with  $pCO_2$  in the whole data and in the subsets, but in

the whole dataset and in the subset FIELD, the correlation coefficients between total nutrients (TN, TP) and  $pCO_2$  were higher than between TOC and  $pCO_2$  (Table 4). The slopes of regression equations between TOC and  $pCO_2$  in the subsets varied. For example with the same concentration of TOC,  $pCO_2$  was higher in the subset FIELD than in the subset PEAT (Fig. 6b). If the subset MINERAL had been plotted in Fig. 6b, the observations would have been situated in between the subsets FIELD and PEAT.

Also, the catchment characteristics had significant correlations with  $CO_2$ . The best predictor was agricultural land in the catchment, which had a significant positive correlation with  $pCO_2$  in the whole dataset and in the subsets FIELD and MIN-ERAL (Table 4). Built-up land in the catchment also increased  $pCO_2$  in lakes. The proportion of water in the catchment had a weak negative relationship with  $pCO_2$  in many subsets (Table 4).

### Discussion

Sources of total inorganic carbon

Weathering in the catchments appeared to dominate the input of TIC into lakes. The main land use types that had a positive correlation with TIC in lakes were agricultural and built-up land in the catchment. There was a strong relationship between TIC and  $CO_2$  in



Table 4 Pearson
correlation coefficients
between $log[pCO_2]$ and
catchment and water quality
variables for the whole
dataset and for the subsets
FIELD, PEAT and
MINERAL

Whole data **FIELD PEAT** MINERAL 790 120 117 119 -0.208\*\*\*NS NS NS Peatland (%) Agricultural land (%) 0.403\*\*\* 0.325\*\*\* NS 0.398\*\* 0.329\*\*\* 0.337\*\*\* Built-up (%) NS NS Water (%) -0.217\*\*\*-0.313\*\*-0.286\*\*NS Forest (%) NS NS 0.202\*NS Latitude -0.402\*\*\*NS NS -0.615\*\*\* $TOC (mg l^{-1})$ 0.396\*\*\* 0.367\*\*\* 0.288\*\* 0.698\*\*\* TN ( $\mu g l^{-1}$ ) 0.463\*\*\* 0.681\*\*\* 0.492\*\*\* NS TON ( $\mu g l^{-1}$ ) 0.410\*\*\* 0.245\*NS 0.670\*\*\*  $NH_4 (\mu g l^{-1})$ 0.420\*\*\* 0.296\*\* 0.254\*\* 0.487\*\*\*  $NO_3 (\mu g l^{-1})$ 0.300\*\*\* NS NS 0.300\*\*\* TP ( $\mu g l^{-1}$ ) 0.429\*\*\* 0.432\*\*\* 0.623\*\*\* 0.217\*TIC (mg  $l^{-1}$ ) 0.649\*\*\* 0.445\*\*\* 0.734\*\*\* 0.461\*\*\* Ca  $(\text{mg l}^{-1})$ 0.599\*\*\* 0.318\*\* 0.600\*\*\* 0.551\*\*\*  $Mg (mg l^{-1})$ 0.559\*\*\* 0.383\*\*\* 0.540\*\*\* 0.474\*\*\*  $SO_4 (mg l^{-1})$ 0.436\*\*\* NS NS 0.506\*\*\*  $SiO_2 (mg l^{-1})$ 0.344\*\*\* 0.231\* 0.605\*\*\* 0.230\*

See Table 1 for explanation of abbreviations Significant coefficients \*P < 0.05, \*\*P < 0.001, \*\*\*P < 0.0001 are shown, NS not significant

Table 5 Stepwise linear multiple regression equations for  $log[pCO_2]$  for the whole dataset and for the subsets FIELD, PEAT and MINERAL

Analyzed set	Parameters							
	α	log[Ca]	log[TOC]	log[TP]	log[TN]	log[SiO <sub>2</sub> ]	$R^2$	
Whole data	5.78	0.353 <sup>a</sup>		0.230 <sup>b</sup>		0.103°	0.53	
FIELD	4.11	0.223 <sup>e</sup>			0.401 <sup>a</sup>	$0.0998^{b}$	0.34	
PEAT	6.34	0.301 <sup>b</sup>				$0.200^{a}$	0.47	
MINERAL	5.64	$0.302^{b}$	0.444 <sup>a</sup>			0.111 <sup>c</sup>	0.68	

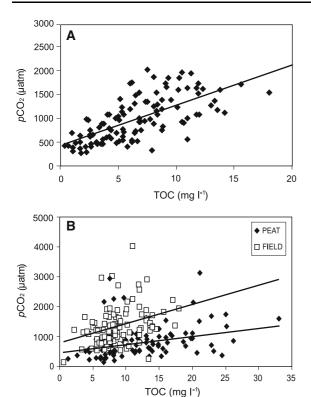
See Table 1 for explanation of abbreviations. The order of independent variables selected by the regression is denoted by letters<sup>a-c</sup>

lakes but no correlation between TIC and TOC, suggesting that decomposition of organic matter was generally not an important source for the main part of TIC, although decomposition contributed to the  $\rm CO_2$  content of the lakes. Billett et al. (2007) proposed different sources for DOC and  $\rm CO_2$  in UK peatland streams on the basis of the different ages and  $\delta^{13}$  signatures of DOC and  $\rm CO_2$ . The main part of  $\rm HCO_3^-$  in natural waters of Finland has been proposed to originate from biogenic carbon acid taking part in weathering (Lahermo et al. 1996). This is supported by our results, because TIC in lakes correlated with Ca and Mg, which also originate mainly from weathering. TIC was better predicted by catchment characteristics than  $p\rm CO_2$ , further

suggesting that although TIC mainly enters the lakes from the surrounding catchments, CO<sub>2</sub> is also produced within the lake ecosystem.

Lakes in agricultural areas had higher TIC concentrations than lakes in peatland rich areas. Raymond and Cole (2003) obtained similar results in Mississippi river basin, where the alkalinity export increased as a function of the percentage of cropland in the catchment, whereas forests in the catchment decreased the alkalinity export. Agriculture is concentrated on fine grained soils, where weathering reactions are more efficient due to greater specific surface area. Furthermore, cultivation of these soils enhances weathering by increasing the interaction with water. Correspondingly, in a large northern river





**Fig. 6** The relationship between **a** the TOC content (mg l<sup>-1</sup>) of the lake water and  $pCO_2$  ( $\mu$ atm) in the subset MINERAL (y = 85.0x + 420;  $R^2 = 0.49$ ). **b** The TOC content (mg l<sup>-1</sup>) of the lake water and  $pCO_2$  ( $\mu$ atm) in the subset PEAT, (y = 25.7x + 577;  $R^2 = 0.08$ ) and in the subset FIELD, (y = 64.5x + 777;  $R^2 = 0.13$ )

basin, Ottawa River, soil respiration and carbonate weathering were proposed to be the main sources of DIC, whereas in-river respiration and photosynthesis were not significant in the river carbon budgets. The highest DIC in Ottawa River was recorded in a subbasin rich in carbonates and heavily used for agriculture (Telmer and Veizer 1999).

In the PCA and correlation analysis, agricultural land and built-up land were suggested to be important catchment sources of TIC in the whole dataset and in the subset FIELD. In the subset PEAT, forests on mineral soil were positively associated with TIC suggesting that mineral soils are more important sources of TIC than peatlands.

# Sources of TOC in different land uses

Peatlands were proposed to be important catchment sources of TOC in the whole dataset and in the subset PEAT. However, also agricultural land in the catchment had a positive relationship with TOC concentrations in the subsets MINERAL and PEAT. TOC concentrations were relatively high in both FIELD and PEAT, but in the subset of PEAT, water color was considerably higher than in the subset FIELD, partly due to higher Fe concentrations in PEAT. In addition, the average TOC:TON-ratio was significantly lower (19 vs. 30) in the subset FIELD compared to the subset PEAT, suggesting a different origin of TOC in the lakes situated in agricultural areas compared to the lakes in peatland dominated areas. DOC of algal origin has little color, compared to the colored humic substances of terrestrial DOC (Meili 1992) and furthermore, terrestrially fixed organic matter typically has a higher TOC:TON ratio compared to phytoplankton. The low C:N rations close to the Redfield TOC:TON ratio of 6.6:1 indicate algal-derived material (Redfield et al. 1963). On the basis of the whole-lake carbon isotope additions, Bade et al. (2007) showed that algal contribution to the DOC pool was 40% in the nutrient enriched lake and 5% in the more humic lake. In humic lakes the dark water color may also limit primary production.

Finnish lakes are shallow and extended littoral zones are typical. Littoral zones are important producers of organic carbon in the lake ecosystem (Wetzel 1983), and their influence on the whole-lake C budget should not be underestimated. Lauster et al. (2006) suggested that littoral zones increase whole-lake net ecosystem production especially in eutrophic systems. In our data, the ratio of shoreline length to lake area was the most important predictor for TOC (r = 0.40, P < 0.001) in the subset FIELD. The ratio of shoreline length to the lake area describes the shape of the lake; the greater the ratio, the more sheltered bays there are enabling the formation of dense littoral vegetation. Andersson and Kumblad (2006) found that the pelagial is fed with carbon fixed by primary producers in benthic and littoral zones, and that there was a strong interaction between the habitats.

Principal component analysis and the correlation analysis showed a strong relationship between TOC and TON in the subset MINERAL suggesting a more uniform origin of organic matter compared to the subsets PEAT and FIELD. In the subset MINERAL, the catchments are mainly forests on mineral soil, whereas in the subsets PEAT and FIELD the land use is more variable.



Sources of CO<sub>2</sub> during the autumn high flow period

Two sources of CO<sub>2</sub> in lakes could be identified by statistical analysis of the data: weathering processes in the catchment, supported by the correlation between CO<sub>2</sub> and Ca, Mg, SiO<sub>2</sub> and TIC, and the decomposition of organic matter, supported by the correlation between CO<sub>2</sub> and TOC, TN and TON. In solute budget constructed for the Ipswich River basin, Williams et al. (2005) found that chemical weathering was the largest source of Ca, Mg and SiO<sub>2</sub>. In their study, and in our data, the agricultural and urban areas of the catchment had strong positive correlations with base cation concentrations in the water. Billett et al. (2007) suggested that carbonate weathering may be an important source of CO<sub>2</sub> in UK peatland streams. The importance of the weathering processes in the catchment as a source of TIC and CO<sub>2</sub> may be overemphasized in our data due to autumn sampling. In autumn, the water discharge from the catchments is higher compared to winter and summer, and therefore, in-lake processes are expected to dominate in winter and summer, whereas the influence of the catchments is highlighted in autumn. The importance of decomposition processes as a source of CO<sub>2</sub> in Finnish lakes during stagnation periods was shown in Striegl et al. (2001) and Kortelainen et al. (2006).

Lake water  $CO_2$  correlated with TOC suggesting that decomposition of organic matter is another important source of  $CO_2$  in lakes as shown in Swedish lakes by Sobek et al. (2003). In our study,  $CO_2$  was also strongly associated with TN, TON,  $NO_3$ ,  $NH_4$  and TP, implying the importance of the quality of organic matter and availability of nutrients for the decomposition processes.

In Finnish lakes most of the TN is organic (88% in the whole dataset). A loose connection between TOC and TON in the whole dataset and in the subsets PEAT and FIELD suggests several sources of organic matter. In the subset FIELD, a low TOC:TON ratio indicates an algal contribution. Raymond and Bauer (2001) showed that heterotrophic bacteria in the river water preferred young labile DOC over old refractory DOC, and furthermore, autochthonous carbon was preferentially utilized by bacteria compared to terrestrial DOC (Kritzberg et al. 2004). Bacterial growth on allochthonous organic carbon in lake water was also shown to be more efficient after addition of

inorganic nutrient (N and P) (Reche et al. 1998). The strong negative correlation between percentage of water in the catchment and TOC concentrations indicates higher retention (either degradation or accumulation) of organic matter in lakes in the subset FIELD compared to the other subsets. Furthermore, the water soluble and bioavailable organic carbon have been shown to be higher in agricultural soils than in forest soils (Boyer and Groffman 1996).

Higher CO<sub>2</sub> as well as TIC in lakes in agricultural areas compared to other subgroups corroborates the hypothesis of the multiple sources of CO<sub>2</sub>. The high availability of organic matter for decomposition in lakes contributes to high  $pCO_2$  in the subset FIELD, but weathering is also more efficient in disturbed, finestructured soils. Dark water color in the subset PEAT may also restrict the photochemical decomposition of TOC in the subsurface layers, and consequently, only a part of TOC may by available for further decomposition (Granéli et al. 1996). Part of the difference in inorganic carbon concentrations between the subsets PEAT and FIELD is due to the lower equilibrium level of lake water CO<sub>2</sub> in the subset PEAT, because lakes situated in catchments with large areas of peatland are typically acidic due to organic acidity (e.g. Kortelainen and Mannio 1988; Kortelainen et al. 1989), whereas lakes in agricultural areas are more alkaline due to chemical weathering and regular liming.

The domination of decomposition processes versus weathering processes as a principal regulator of CO<sub>2</sub> concentrations in lakes varied between the subgroups, as suggested by the PCA analysis. In the whole dataset with mixed land use the dual sources of  $pCO_2$ were emphasized, whereas different processes dominated pCO<sub>2</sub> concentrations in the subsets. In the subset PEAT, where on average 48% of each catchment consists of forests on mineral soils (Table 1), weathering in the catchments controlled CO<sub>2</sub> probably due to the refractory nature of organic matter and the unfavorable conditions for decomposition processes (low nutrient availability and dark water restricting photo-oxidation). In the subset FIELD, decomposition mainly controlled CO<sub>2</sub>, supported by the close connection with TON, NO<sub>3</sub> and NH<sub>4</sub> in the PCA analysis. Both processes were indicated to be important in the subset MINERAL. The results also indicated that more uniform is the land use in the catchment, and consequently the quality of the allochthonous organic matter, the



stronger is the connection between TOC and  $CO_2$ . This is supported by the strong relationship in PCA and correlation analysis between lake water TOC and  $pCO_2$  in the subset MINERAL, with the most uniform land use of all the subsets.

Silicate (SiO<sub>2</sub>) correlated positively with both TIC and CO<sub>2</sub> and was included in the multiple regression models for CO<sub>2</sub>. The decomposition of organic matter produces organic acids and carbon dioxide, both of which enhance weathering and thus SiO<sub>2</sub> concentrations. The increase in the proportion of vegetation cover increased the fluxes of weathering products such as dissolved silica in arctic Swedish rivers (Humborg et al. 2004; Smedberg et al. 2006) and furthermore, there was a positive relationship between TOC and dissolved silica in arctic rivers. For example some wetland plants have been shown to accumulate large quantities of amorphous silica (Struyf and Conley 2008). Probably due to this dual relationship, in the PCA analysis of the whole dataset SiO<sub>2</sub> is not associated with either of the two main principal components and has little influence on  $pCO_2$  in lakes.

#### **Conclusions**

Organic carbon dominated the TC pool in boreal Finnish lakes. The proportion of inorganic carbon was highest in agricultural areas. Weathering was the main source of TIC in lakes, and in agricultural areas both weathering and export of inorganic carbon from the catchments was enhanced by fine structure of soils. Furthermore, due to the high pH of these lakes, a smaller part of TIC was in gaseous form.

Two contrasting sources of  $CO_2$  in lakes could be identified in the statistical analysis of the data: weathering processes in the catchments and decomposition of organic matter were both important sources of  $CO_2$  in boreal, Finnish lakes during autumn overturn. The importance of these two processes as the regulators of  $CO_2$  in lakes varied in different catchment types. In agricultural areas and in the catchments of forested mineral soils, decomposition processes were suggested to be the main regulator of  $CO_2$ , whereas in peatland rich areas (on average 44% peatlands), weathering processes controlled  $CO_2$ .  $CO_2$  was also strongly associated with total nutrients TN and TP, implying the importance of the quality of organic matter for decomposition. Highest  $CO_2$  in lakes was recorded in

agricultural areas: weathering is more efficient in disturbed, fine-structured soils, and furthermore, high availability of nutrients in these lakes enhanced both primary production and decomposition.

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## References

- Andersson E, Kumblad L (2006) A carbon budget for an oligotrophic clearwater lake in mid-Sweden. Aquat Sci 68:52–64
- Bade DL, Carpenter SR, Cole JJ et al (2007) Sources and fates of dissolved organic carbon in lakes as determined by whole-lake carbon isotope additions. Biogeochemistry 84:115–129
- Billett MF, Garnett MH, Harvey F (2007) UK peatland streams release old carbon dioxide to the atmosphere and young dissolved organic carbon to rivers. Geophys Res Lett 34:L23401
- Boyer JN, Groffman PM (1996) Bioavailability of water extractable organic carbon fractions in forest and agricultural soil profiles. Soil Biol Biochem 28(6):783–790
- Butler JN (1982) Carbon dioxide equilibria and their applications. Addison-Wessley Publishing Company, Reading
- Canham CD, Pace MP, Papaik MJ et al (2004) A spatially explicit watershed-scale analysis of dissolved organic carbon in Adirondack lakes. Ecol Appl 14:839–854
- Cole JJ, Prairie YT, Caraco NF et al (2007) Plumbing the global carbon cycle: integrating inland waters into the terrestrial carbon budget. Ecosystems 10:171–184
- Dawson JJC, Billett MF, Neal C et al (2002) A comparison of particulate, dissolved and gaseous carbon in two contrasting upland streams in the UK. J Hydrol 257:226–246
- Dawson JJC, Billett MF, Hope D et al (2004) Sources and sinks of aquatic carbon in a peatland stream continuum. Biogeochemistry 70:71–92
- Dillon PJ, Molot LA (1997) Dissolved organic and inorganic carbon mass balances in central Ontario lakes. Biogeochemistry 36:29–42
- Findlay SEG, Sinsabaugh RL (2003) Aquatic ecosystems: interactivity of dissolved organic matter. Academic Press, San Diego
- Findlay SEG, Quinn JM, Hickey CW (2001) Effects of land use and riparian flowpath on delivery of dissolved organic carbon to streams. Limnol Oceanogr 46(2):345–355
- Gaillardet J, Dupré B, Louvat P et al (1999) Global silicate weathering and CO<sub>2</sub> consumption rates deduced from the chemistry of large rivers. Chem Geol 159:3–30
- Granéli W, Lindell M, Tranvik L (1996) Photo-oxidative production of dissolved inorganic carbon in lakes of different humic content. Limnol Oceanogr 41(4):698–706
- Henriksen A, Skjelkvåle BL, Lien L et al (1996) Regional lake surveys in Finland–Norway–Sweden–Northern Kola–Russian Karelia–Scotland–Wales 1995. Coordination and design. Acid rain research report 40, NIVA, Oslo



- Humborg C, Smedberg E, Blomqvist S et al (2004) Nutrient variations in boreal and subarctic Swedish rivers: landscape control of land–sea fluxes. Limnol Oceanogr 49(5):1871–1883
- IPCC (2001) In: Houghton JT, Ding Y, Griggs DJ et al (eds) Climate change 2001: the scientific basis. Contribution of working group I to the third assessment report of the intergovernmental panel on climate change. Cambridge University Press, Cambridge
- Johnson MS, Lehmann J, Couto EG et al (2006) DOC and DIC in flowpaths of Amazonian headwater catchments with hydrologically contrasting soils. Biogeochemistry 81: 45–57
- Juutinen S, Rantakari M, Kortelainen P et al (2009) Methane dynamics in different boreal lake types. Biogeosciences (in press)
- Kling GW, Kipphut GW, Miller MC (1992) The flux of CO<sub>2</sub> and CH<sub>4</sub> from lakes and rivers in arctic Alaska. Hydrobiologia 240:23–46
- Kortelainen P, Mannio J (1988) Natural and anthropogenic acidity sources for Finnish lakes. Water Air Soil Pollut 42:341–352
- Kortelainen P, Mannio J, Forsius M et al (1989) Finnish lake survey: the role of organic and anthropogenic acidity. Water Air Soil Pollut 46:235–249
- Kortelainen P, Rantakari M, Huttunen JT et al (2006) Sediment respiration and lake trophic state important predictors for the large CO<sub>2</sub> evasion from small boreal lakes. Glob Chang Biol 12:1554–1567
- Kritzberg ES, Cole JJ, Pace ML (2004) Autochthonous versus Allochthonous carbon sources of bacteria: results from whole-lake <sup>13</sup>C addition experiments. Limnol Oceanogr 49(2):588–596
- Lahermo P, Väänänen P, Tarvainen T et al (1996) Suomen geokemian atlas. Osa 3: Ympäristögeokemia–purovedet ja-sedimentit (Geochemical atlas of Finland. Part 3: environmental geochemistry—stream waters and sediments). Geologian tutkimuskeskus, Espoo. Geological survey of Finland, Espoo
- Lauster GH, Hanson PC, Kratz TK (2006) Gross primary production and respiration differences among littoral and pelagic habitats in northern Wisconsin lakes. Can J Fish Aquat Sci 63:1130–1141
- Mannio J, Räike A, Vuorenmaa J (2000) Finnish lake survey 1995: regional characteristics of lake chemistry. Verh Internat Verein Limnol 27:362–367
- Meili M (1992) Sources, concentrations and characteristics of organic matter in softwater lakes and streams of the Swedish forest region. Hydrobiologia 229:23–41
- Neal C, Hill S (1994) Dissolved inorganic and organic carbon in moorland and forest streams: Plynlimon, mid Wales. J Hydrol 153:231–243
- Palmer SM, Hope D, Billett M, Dawson JJC, Bryant CL (2001) Sources of organic and inorganic carbon in a headwater stream: evidence from carbon isotope studies. Biogeochemistry 52:321–338
- Plummer LN, Busenberg E (1982) The solubility of calcite, aragonite, and vaterite in CO<sub>2</sub>–H<sub>2</sub>O solutions between 0 and 90°C, and an evaluation of the aqueous model for CaCO<sub>3</sub>–CO<sub>2</sub>–H<sub>2</sub>O equilibria. Geochim Cosmochim acta 46:1011–1040
- Rantakari M, Kortelainen P, Vuorenmaa J et al (2004) Finnish lake survey: the role of catchment attributes in

- determining nitrogen, phosphorus and organic carbon concentrations. Water Air Soil Pollut Focus 4:683–699
- Raymond PA, Bauer JE (2001) Riverine export of aged terrestrial organic matter to the North Atlantic Ocean. Nature 409:497–500
- Raymond PA, Cole JJ (2003) Increase in the export of alkalinity from North America's largest river. Science 301:88–91
- Raymond PA, Oh N-H, Turner RE, Broussard W (2008) Anthropogenically enhanced fluxes of water and carbon from the Mississippi river. Nature 451:449–452
- Reche I, Pace ML, Cole JJ (1998) Interactions of photobleaching and inorganic nutrients in determining bacterial growth on colored dissolved organic carbon. Microb Ecol 36:270–280
- Redfield AC, Ketchum BH, Richards FA (1963) The influence of organisms on the composition of seawater. In: Hill MN (ed) The sea, vol 2. Wiley, Chicherster
- Salonen K, Vähätalo A (1994) Photochemical mineralisation of dissolved organic matter in lake Skjervatjern. Environ Int 20(3):307–312
- SAS Institute (2001) SAS version 8.2 for windows. SAS Institute Inc, Cary
- Sinsabaugh RL, Findley SEG (2003) Dissolved organic matter: out of the black box into the main stream. In: Findlay SEG, Sinsabaugh RL (eds) Aquatic ecosystems: interactivity of dissolved organic matter. Academic Press, San Diego
- Smedberg E, Mörth C-M, Swaney DP et al (2006) Modeling hydrology and silicon–carbon interactions in taiga and tundra biomes from a landscape perspective: implications for global warming feedbacks. Glob Biogeochem Cycles 20:GB2014
- Sobek S, Algesten G, Bergström A-K et al (2003) The catchment and climate regulation of *p*CO<sub>2</sub> in boreal lakes. Glob Chang Biol 9:630–641
- Striegl RG, Kortelainen P, Chanton JP et al (2001) Carbon dioxide partial pressure and <sup>13</sup>C content of north temperate and boreal lakes at spring ice melt. Limnol Oceanogr 46:941–945
- Striegl RG, Dornblaser MM, Aiken GR et al (2007) Carbon export and cycling by the Yukon, Tanana, and Porcupine rivers, Alaska, 2001–2005. Water Resour Res 43:W02411
- Struyf E, Conley DJ (2008) Silica: an essential nutrient in wetland biogeochemistry. Front Ecol Environ 6. DOI 10.1890/070126
- Stumm W, Morgan JJ (1981) Aquatic chemistry. An introduction emphasizing chemical equilibria in natural waters. Wiley, New York
- Telmer K, Veizer J (1999) Carbon fluxes, pCO<sub>2</sub> and substrate weathering in a large northern river basin, Canada: carbon isotope perspectives. Chem Geol 159:61–86
- Walvoord MA, Striegl RG (2007) Increased groundwater to stream discharge from permafrost thawing in the Yukon River basin: potential impacts on lateral export of carbon and nitrogen. Geophys Res Lett 34:L12402
- Wetzel RG (1983) Limnology, 2nd edn. Saunders Collage Publishing, Philadelphia
- Williams M, Hopkinson C, Rastetter E et al (2005) Relationships of land use and stream solute concentrations in the Ipswich river basin, northeastern Massachusetts. Water Air Soil Pollut 161:55–74

